

1. Introduction

On 19 May 2013, two supercells moved through central Oklahoma and produced seven tornadoes, including one of EF-3 and one of EF-4 intensity. High temporal and spatial resolution observations were collected in both strong tornadoes by a mobile, rapid-scan, X-band, polarimetric radar (RaXPol). In conjunction with the nearby WSR-88D Twin Lakes radar, dual-Doppler analyses were synthesized to produce the three-dimensional wind field and analyses of vertical vorticity throughout the tornadoes' lifecycles. Presented



herein is a vertical vorticity analysis of the Norman-Shawnee supercell, which includes data from three separate RaXPol deployments (left). Analysis of the dual-Doppler windfield with respect to differential

reflectivity (Z_{DR}) columns are also examined for the Edmond-Carney supercell (top right).

2. Methods and Radar Characteristics Characteristics of the radars used to collect data on 19 May 2013, as well as the methods used to process and create the dual-Doppler wind syntheses,

are described below:



	RaXPol	RADAR
	20 kW	Peak Tr
	9.73 GHz	Operat
	1.0°	Beamw
	75 m	Range F
	180° s ⁻¹	Antenn

High-temporal resolution -- radar scans every 2 seconds

Analy

RaXPol



Objective Analysis

Observation Processing and Wind Syntheses (OPAWS) (e.g., Kosiba et al. 2013) was used to map the data to a Cartesian grid using a two-pass Barnes analysis, and a grid spacing of 250 m.

KTLX Analyses

 Spatially variable advection correction technique employed to interpolate KTLX data linearly, in a Lagrangian sense, to the RaXPol analysis times (Shapiro et al. 2010a, b).

 Interpolated KTLX volumes are paired with RaXPol volumes at each corresponding time to create a dual-Doppler synthesis valid at each RaXPol volume.

Dual-Doppler Wind Synthesis Z_D Column Analysis

DUAL-DOPPLER AND POLARIMETRIC OBSERVATIONS OF TWO TORNADIC SUPERCELLS IN CENTRAL OKLAHOMA ON 19 MAY 2013 Zachary B. Wienhoff¹, Howard B. Bluestein¹, Jana B. Houser², Jeffrey C. Snyder³, Alan Shapiro¹, Corey K. Potvin^{3, 4}

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ABOVE: Evolution of the low-level vertical vorticity (ζ) at six selected times (labeled above). Contours of ζ are in increments of 0.5x10⁻² s⁻¹. Black contours represent cyclonic vorticity ($\zeta > 0$), and blue contours represent anticyclonic vorticity ($\zeta < 0$). Horizontal wind vectors are storm relative.





4. Z_{DR} Column Analysis - Carney-Edmond Supercell 21:51:43 UTC 21:48:15 UTC 0 X distance from RaXPol X distance from RaXPol

5. Conclusions

This study demonstrates that it is possible to use time interpolated radar data to supplement rapid-scan radar data where a second radar can be used to provide additional insight on the physical processes within a convective storm. Qualitatively, this technique yielded reasonable results. However, additional studies are needed to better understand the quantitative results (e.g., how well are the updraft strengths represented) and identify areas where this technique may exacerbate dual-Doppler errors. Acknowledgments

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